

Natural and man-made gamma emitters in Gulf of Eilat / Aqaba sediments

**Daniela Pittauerová¹ Gerald Kirchner²
Dieter Garbe-Schönberg³ Ami Nishri⁴ Barak Herut⁴
Helmut W. Fischer¹**

¹University of Bremen, Institute for Environmental Physics, Germany

²Bundesamt für Strahlenschutz, Berlin, Germany

³Israel Oceanographic and Limnological Research, Haifa, Israel

⁴Christian-Albrechts-University of Kiel, Institute of Geosciences, Kiel, Germany

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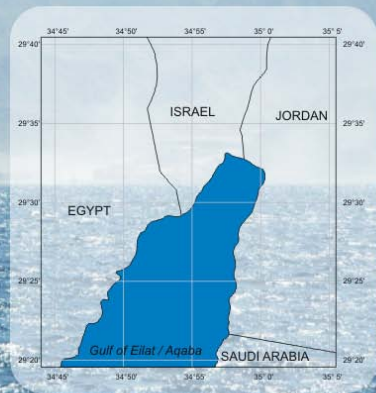
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- Pb-210 and Cs-137 sedimentation rates
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Gulf of Eilat



Motivation

- Last century's accelerated anthropogenic pollution and input of nutrients → negative effect on marine environment in Gulf of Eilat / Aqaba (GOE) - coral reefs and marine life biodiversity.
- Gamma emitting radionuclides in sediment cores analyzed within a study of sources and effects of particulate phosphorous in GOE.
- Particulate phosphorous sources: mariculture, sewage and phosphate ore dust from industrial ports in Aqaba and Eilat.
- Estimated P release from the port of Eilat: $\geq 8 \cdot 10^6$ mol P, Aqaba port approximately 10-fold higher.
- No previous publications on radionuclides in sediment profiles, sedimentation rates and radionuclide inventories from the studied area.

Sampling: sediments

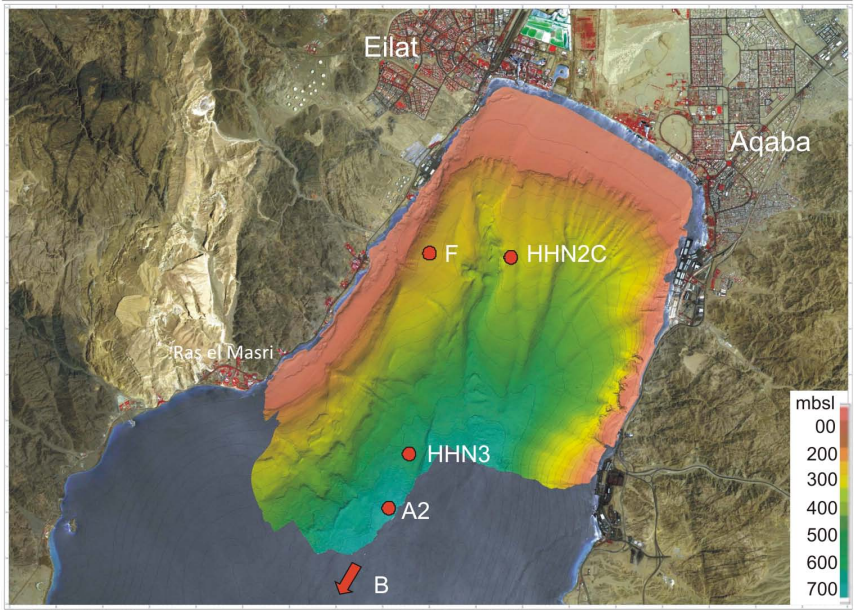
- Five short sediment cores taken during 2007-2008
- St. F and HHN2C: shallow (240-316 m)
- St. A2 and HHN3: deeper part of Eilat subbasin (600-700 m)
- St. B: further south in Eilat deep (800 m), a reduced effect of anthropogenic pollution expected



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Measurements

Gamma spectrometry

- Samples hermetically sealed - waited for equilibrium for ^{226}Ra determination
- Low-level low-background γ -spec., 50% HPGe coaxial detector
- LabSOCS for a characterized detector used for efficiency calculations - variable geometries
- Cascade summing corrections applied
- Samples: 1-10 g
- Counting times: 2-3 days for small samples, 1-2 days for larger samples
- Gamma emitters: ^{210}Pb , ^{226}Ra (^{214}Pb , ^{214}Bi), ^{40}K , ^{228}Ra (^{228}Ac), ^{228}Th (^{212}Pb , ^{208}Tl), ^{137}Cs

Age models

$^{210}\text{Pb}_{\text{xs}}$

- Constant flux - constant sedimentation (CF-CS) model
- Variations in depth profile likely to be caused by other factors than changes of sedimentation rate

Age models

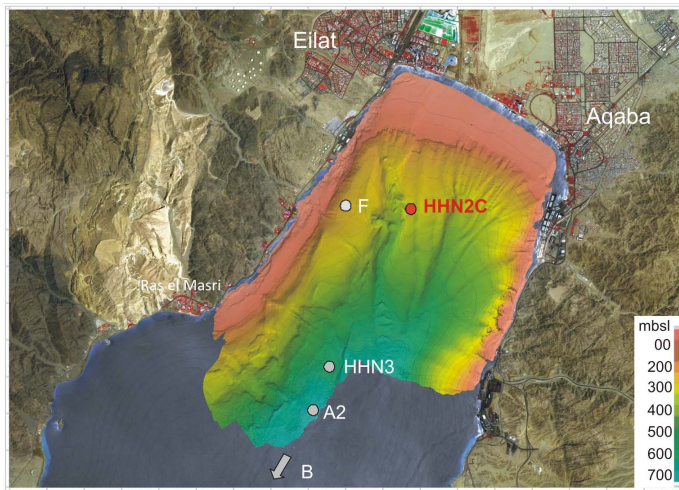
$^{210}\text{Pb}_{xs}$

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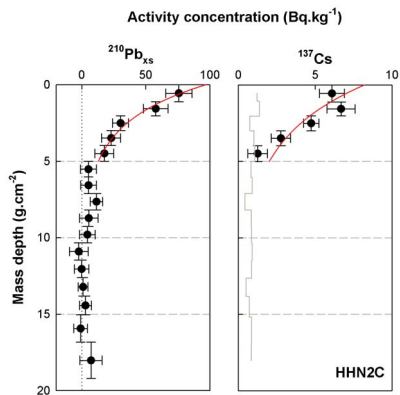
^{137}Cs

- Additional tracer
- Main source: global weapon test fallout, minimal effect of Chernobyl fallout
- No clear 1963 maxima in profiles
- Instead, “exponential” decay in some profiles: attempt to use the same CF-CS model

Shallow core HHN2C



Shallow core HHN2C



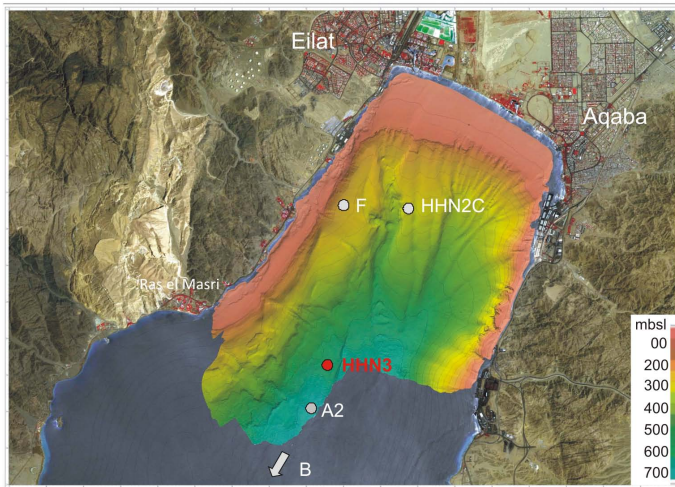
$$r(^{210}\text{Pb}_{\text{xs}})$$

$$0.076 \pm 0.008 \text{ g} \cdot \text{cm}^{-2} \text{ yr}^{-1}$$

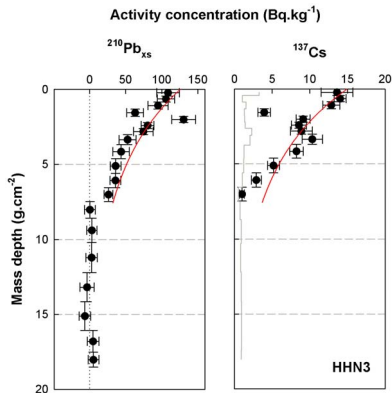
$$r(^{137}\text{Cs})$$

$$0.082 \pm 0.030 \text{ g} \cdot \text{cm}^{-2} \text{ yr}^{-1}$$

Deep core HHN3



Deep core HHN3



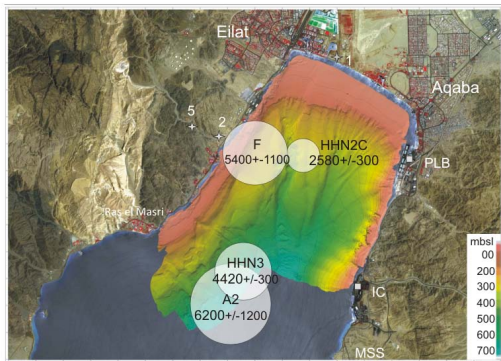
$$r (^{210}\text{Pb}_{\text{xs}})$$

$$0.16 \pm 0.04 \text{ g} \cdot \text{cm}^{-2} \text{ yr}^{-1}$$

$$r (^{137}\text{Cs})$$

$$0.11 \pm 0.03 \text{ g} \cdot \text{cm}^{-2} \text{ yr}^{-1}$$

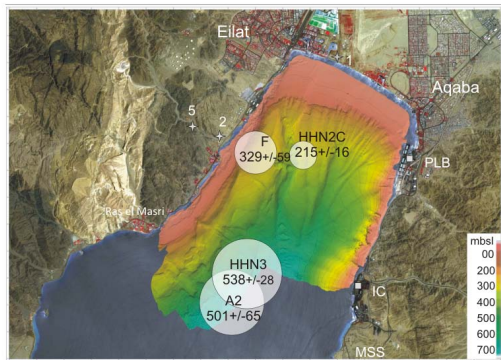
Inventories of $^{210}\text{Pb}_{\text{xs}}$ ($\text{Bq} \cdot \text{m}^{-2}$)



Core	$^{210}\text{Pb}_{\text{xs}}$	
	Inventory $\text{Bq} \cdot \text{m}^{-2}$	Flux $\text{Bq} \cdot \text{m}^{-2} \text{yr}^{-1}$
F	5400 ± 1100	167 ± 33
HHN2C	2580 ± 300	80.6 ± 9.5
HHN3	4420 ± 300	138.1 ± 9.5
A2	6200 ± 1200	195 ± 35
B	5090 ± 310	–

A mean atmospheric flux over continents in latitudal band 10° – 30° N: $160 \text{ Bq} \cdot \text{m}^{-2} \text{yr}^{-1}$ (global compilation, Preiss et al. 1996).

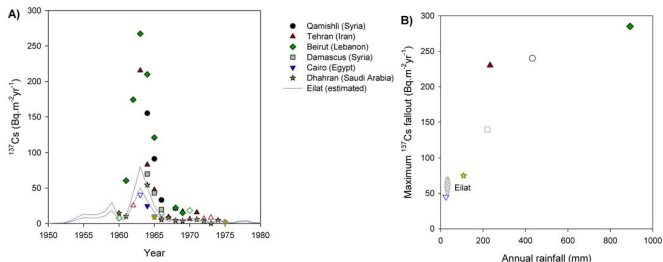
Inventories of ^{137}Cs ($\text{Bq} \cdot \text{m}^{-2}$)



Core	Inventory $\text{Bq} \cdot \text{m}^{-2}$
F	$\geq 329 \pm 59$
HHN2C	215 ± 16
HHN3	538 ± 28
A2	501 ± 65
B	$\geq 400 \pm 26$

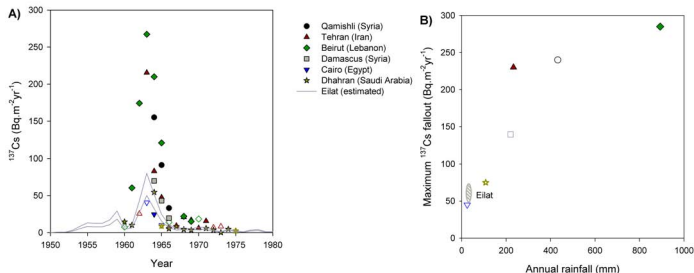
ADDITIONAL: Global weapon test fallout, Middle East: ^{137}Cs

- Data: Environmental Measurements Laboratory Global Fallout Deposition program (on-line database). The measurement series not continuous (full / empty symbols).
- $^{90}\text{Sr} \rightarrow ^{137}\text{Cs}$: constant ratio $\text{Cs}/\text{Sr}=1.5$ assumed
- Solid line: an estimate of fallout in GOE based on UNSCEAR (2000) deposition history scaled to maximal yearly fallout estimate.



ADDITIONAL: Global weapon test fallout, Middle East: ^{137}Cs

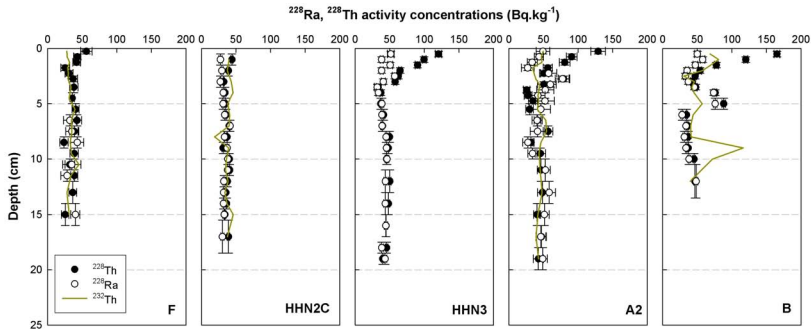
- Fallout varies with latitude and rainfall.
- Maximum ^{137}Cs yearly fallout (in 1963) vs. annual rainfall.
Full symbols represent actual measured values, empty symbols extrapolation when 1963 value was not available.
- The maximum (1963) at GOE estimated $50\text{--}75 \text{ Bq} \cdot \text{m}^{-2} \text{yr}^{-1}$.
- Total fallout: $240\text{--}360 \text{ Bq} \cdot \text{m}^{-2}$ (decay corrected to 2007).



Global weapon test fallout, Middle East: ^{137}Cs

- Compilation of data: Environmental Measurements Laboratory Global Fallout Deposition program (on-line database) - 6 stations in Syria, Iran, Lebanon, Egypt and Saudi Arabia.
- Fallout varies with latitude and rainfall.
- Total fallout in Eilat: $240\text{--}360 \text{ Bq} \cdot \text{m}^{-2}$ (decay corrected to 2007).
- Inventories measured within the study: up to $540 \text{ Bq} \cdot \text{m}^{-2}$
- 33–56% ^{137}Cs not deposited directly, rather erosion derived

Th series radionuclides

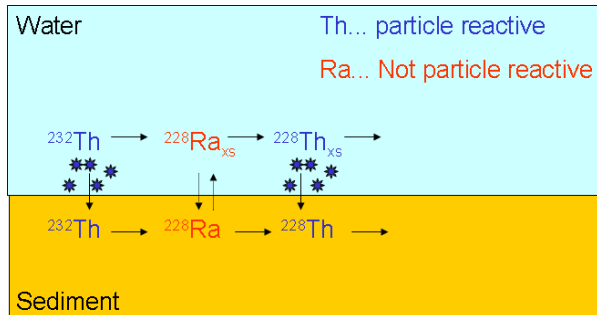


- Gamma emitters: ^{228}Ra and ^{228}Th
- ^{232}Th measured by ICP-MS

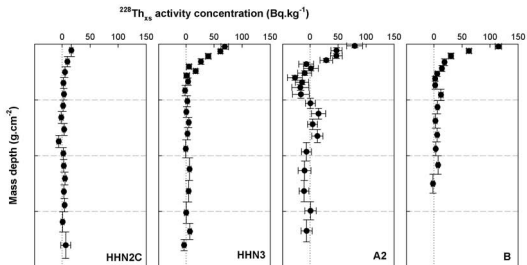
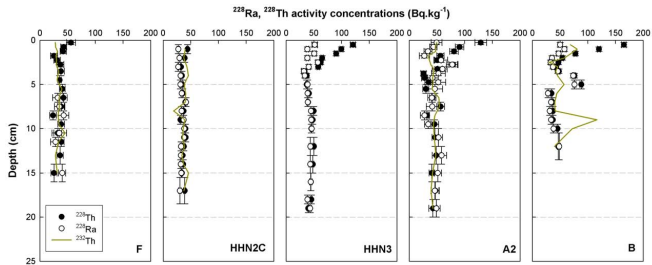
Th series radionuclides

$$^{228}\text{Ra}(t) = ^{232}\text{Th}(0) \cdot (1 - e^{-\lambda_2 t}) + ^{228}\text{Ra}(0) \cdot e^{-\lambda_2 t}$$

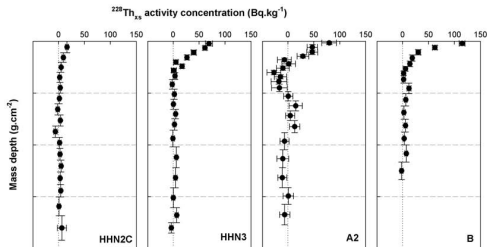
$$^{228}\text{Th}(t) = \lambda_3 \cdot \lambda_2 \cdot ^{232}\text{Th}(0) \cdot \left(\frac{1}{\lambda_2 \cdot \lambda_3} - \frac{e^{-\lambda_2 t}}{\lambda_2(\lambda_3 - \lambda_2)} - \frac{e^{-\lambda_3 t}}{\lambda_3(\lambda_2 - \lambda_3)} \right) + \frac{\lambda_3}{\lambda_3 - \lambda_2} \cdot ^{228}\text{Ra}(0) \cdot (e^{-\lambda_2 t} - e^{-\lambda_3 t}) + ^{228}\text{Th}(0) \cdot e^{-\lambda_3 t}$$



Th series radionuclides - excess ^{228}Th

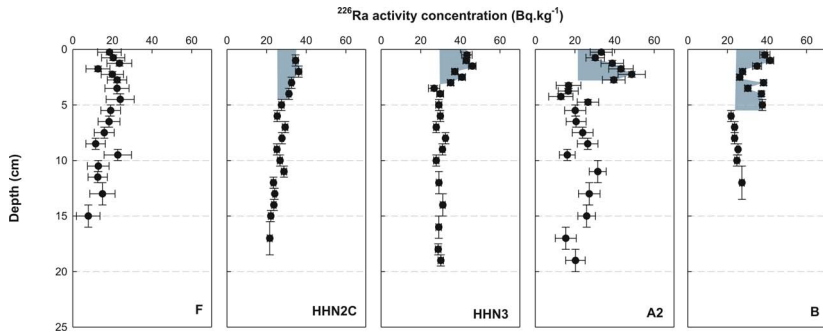


Th series radionuclides - excess ^{228}Th



- Valuable additional information: the core tops were deposited very recently - $T_{1/2}$ (^{228}Th)=1.9 yr.
- Application of a simple CF-CS model leads to several times higher accumulation rates ($0.3\text{--}0.7 \text{ g}\cdot\text{cm}^{-2}\text{yr}^{-1}$) than ^{210}Pb and ^{137}Cs model.
- Reasons: Bioturbation? Ra diffusion? Recent sedimentation rate acceleration? Recent reduction of $^{228}\text{Th}_{xs}$ flux? Postdepositional redistribution?

^{226}Ra - increase in upper parts of the profiles



^{226}Ra "XS" inventories

Core	Top interval cm	μ_{top} $Bq \cdot kg^{-1}$	μ_{bottom} $Bq \cdot kg^{-1}$	t-value	P-value	Inventory of ^{226}Ra "XS" $Bq \cdot m^{-2}$
HHN2C	4.5	33.7	25.4	6.242	0.0004	320 ± 30
HHN3	3.5	40.9	29.5	6.627	0.0004	300 ± 30
A2	3.0	39.0	21.5	5.670	0.0002	410 ± 90
B	5.5	34.8	24.5	5.273	0.0001	440 ± 30

- Increase of ^{226}Ra in the top sections (3–5.5 cm) of 4 cores is statistically significant.
- Phosphate: $1200 Bq \cdot kg^{-1}$ ^{226}Ra
- Estimated phosphate dust release since 1965: $17 \cdot 10^3$ t

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If phosphate dust responsible for ^{226}Ra increase:

- $300\text{--}440 Bq \cdot kg^{-1} \rightarrow 0.25\text{--}0.36 kg \cdot m^{-2}$ of phosphate accumulated on the seabed
- Over the area of $40 km^2 \rightarrow (10.0 - 14.4) \cdot 10^3$ t of phosphate

Summary

Accumulation rates, inventories

- Based on ^{210}Pb and ^{137}Cs CF-CS model: $0.076\text{--}0.22$
 $\text{g} \cdot \text{cm}^{-2} \text{yr}^{-1}$
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- $^{228}\text{Th}_{\text{xs}}$ found in core tops: very fresh sediment
- Common dating model (CF-CS) not applicable

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^{226}Ra

- Increased ^{226}Ra activities in core tops: likely to be caused by contribution of phosphate dust from Eilat and Aqaba industrial ports.
- $(10.0 - 14.4) \cdot 10^3$ t of phosphate material deposited at a seabed of studied area.

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Thank you for your attention!